

Hydrothermal alteration patterns of the Ticlioporphry Cu-Au prospect, Miocene Morocochadistrict, central Peru: surface spatial distribution mineralogy and geochemistry

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### **Abstract**

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The Morococha mining district is part of the Miocene polymetallic belt of central Peru and hosts various types of ore bodies – porphyries, skarns, replacement bodies (mantos) and polymetallic veins. The geology of the district comprises volcanic sequences of the Permian Mito Group, overlain by Triassic to Jurassic carbonates of the Pucará Group, including basalt flow intercalations such as the Montero Basalt, and, mainly in the western part, by Cretaceous carbonate sedimentary sequences. The magmatic activity began at 14.1 Ma with the emplacement of the Anticona diorite intrusion that covers a large area in the western part of Morococha district. During Late Miocene (7-9 Ma) a series of porphyry intrusions (diorites, granodiorites to quartz-monzonites in composition) intruded the different sedimentary sequences as well as the Anticona diorite. Most of these intrusions are barren, but some of them share features with typical porphyry Cu systems. The Ticlio Cu-Au porphyry is an atypical porphyry in the district, where epithermal polymetallic mineralization (Ag, Zn and Pb) is predominant in the peripheral part and where Cu-Mo porphyry systems are dominant in the district such as Toromocho and Codiciada. The Ticlio Cu-Au porphyry is a single granodiorite intrusion that intruded sedimentary layers as well as the Anticona intrusion during the Late Miocene. The Ticlio intrusion is located in the westernmost part of Morococha district. The studied area represents about 2.5 km<sup>2</sup> and field works gave the possibility to map this region. The geology of the area consists of Cretaceous dolomitic limestones, the Ticlio intrusion stock in the central part with xenoliths of basalt on the eastern fringe of the intrusion and the Anticona diorite intrusion that delimits the eastern part of the studied area. Different systems of fracturation in time offset the area. The first system displays a NE-SW orientation whereas the last system of fracturation shows a NW-SE trend. Dike and aplitic dike, which are altered, crosscut the Ticlio intrusion. Based on the crosscutting relationships the aplitic dike could be syn-mineralization. Dike that offsets the Anticona intrusion could be post-mineralization. According to the alteration map, the studied area is overprinted by different zones of alteration. Potassic alteration that is characterized mainly by secondary feldspars, biotites, quartz and magnetite is dominant in the central part of the system where quartz-magnetite veins form the stockwork. In the southern Ticlio intrusion, the density of veining is strong to the SE that is accompanied by increasing in the intensity of the potassic alteration. This is interpreted as the focus of the upwelling fluids. The second stage of the alteration is characterized by chlorite-epidote and the late stage of the alteration is illustrated by the sericitic alteration that occurs on the rim of the Ticlio porphyry and developed about 100'000 years after the emplacement of the Ticlio intrusion. This study also discusses the mineralogy and whole-rock geochemistry on the igneous rocks as well as whole-rock geochemistry and elemental mass balance of the hydrothermal alteration zones within Ticlio porphyry copper-gold deposit. Whole-rock analyses display high-K calc-alkaline affinity and indicate magmatic arc signature with a subducted slab component. Microscope observations and electron microprobe analyses of phenocrysts of plagioclase suggest mixing process in the magma chamber. The rare earth elements are often considered to be immobile, unaffected by the leaching and geochemical redistribution processes that characterize hydrothermal alteration. However, recent studies on the REE contents of hydrothermal fluids and hydrothermally altered rocks have shown that, under particular conditions, REE can be

mobilized by hydrothermal alteration. Chondrite-normalized REE plots of altered rock together with mass balance calculations were used to understand the interactions between fluids and rock and to characterize REE behavior in the different alteration zones. Potassic alteration is accompanied by enrichment in K, Si and depletion in Fe, Sr, Ba. These changes attended replacement of plagioclase and amphibole by K-feldspars and biotite respectively. The copper is introduced most probably during this stage of alteration because of the gain in Cu. Gold could be introduced at the same time. In contrast, phyllic alteration displays strong depletion in Mg, Ca, Na, Fe, Sr and Ba that reflect sericitization of alkali feldspars and destruction of ferromagnesian minerals. The addition of Si is consistent with silicification that is a major feature of phyllic alteration. This hydrothermal alteration is accompanied by strong depletion in heavy rare earth elements (HREE). This latter involves acid leaching during the waning stage of the hydrothermal system. In general, major elements and some minor trace elements decrease from the least altered rock to the late alteration. Moreover, depletion in MREE accompanied by depletion in  $P_2O_5$  indicates that apatite do not resist at the phyllic alteration. The Ticlio porphyry shares similar features of typical porphyry systems. Actually, hydrothermal zoned alteration patterns stockwork and skarn characterize the Ticlio stock. Microscopy observations mass balance calculations and REE patterns indicate a saline and oxidizing hydrothermal fluids with an evolution in its acidity that is typical for the porphyry system. Finally, gold occurs as inclusion within the chalcopyrite presents in the xenoliths of basalt that display higher concentration values of gold and copper compared to the least altered and altered rocks of the Ticlio intrusion. Therefore, this area is not propitious for an eventual prospection due to the low grade of gold and copper and the small volume of the Ticlio porphyry deposit.